Volcanism in a compressional Andean setting: A structural and geochronological study of Tromen volcano (Neuquén province, Argentina)

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[1] We document evidence for growth of an active volcano in a compressional Andean setting. Our data are surface structures and 39Ar-40Ar ages of volcanic products on Tromen volcano. Tromen is an active back-arc volcano in the Andean foothills of Neuquén province, Argentina. Its volcanic products are unconformable upon Mesozoic strata of the Neuquén basin. The volcano straddles a N-S trending pop-up, which formed during E-W shortening. The main underlying structures are eastward verging thrusts. Their traces curve around the eastern foot of the volcano. Minor folds and faults also occur in the volcanic cover of Tromen, as a result of E-W shortening. New 39Ar-40Ar ages for these volcanic rocks are younger than 2.27 ± 0.10 Ma and show that Tromen has been active almost continuously from the late Pliocene to the Holocene. We conclude that volcanism and thrusting have been coeval and that magma must have reached the surface in a tectonic setting of horizontal compression. Our results have wider implications for magmatic processes in such settings. Citation: Galland, O., E. Hallot, P. R. Cobbold, G. Ruffet, and J. de Bremond d’Ars (2007), Volcanism in a compressional Andean setting: A structural and geochronological study of Tromen volcano (Neuquén province, Argentina), Tectonics, 26, TC4010, doi:10.1029/2006TC002011.

1. Introduction

[2] Volcanic activity is especially common in areas of extensional tectonics, and notably so at divergent plate boundaries. In principle, such a setting should induce vertical hydraulic fractures, along which magma rises directly through the upper crust, to the surface [Hubbert and Willis, 1957; Sibson, 2003]. This is borne out by field observations, for example, in the rift system of Iceland [Gudmundsson, 1984; Helgason and Zentilli, 1985].

[3] In contrast, a context of crustal thickening, where the least principal stress is vertical, should impede the rise of magma [Watanabe et al., 1999; Sibson, 2003] and produce horizontal hydraulic fractures [Hubbert and Willis, 1957; Sibson, 2003]. Under these conditions, some authors have argued that volcanic activity is infrequent, if not absent [Glazner, 1991; Hamilton, 1995]. Nevertheless, volcanism and compressional tectonics are coeval in many places on the Earth’s surface, for example in the Andes [Allmendinger et al., 1997; Lamb et al., 1997], and in Japan [Townend and Zoback, 2006].

[4] Because it is intuitive that extension provides room for magma, it is sometimes assumed that such extension is a necessity [Glazner, 1991; Hamilton, 1995]. For example, in the Andes the volcanic arc has been interpreted as a result of local extension [e.g., de Silva, 1989; Hamilton, 1995]. However, Andean volcanoes have been active in a regional context of horizontal compression since 50 Ma or earlier [Allmendinger et al., 1997; Lamb et al., 1997; Kendrick et al., 1999; Kendrick et al., 2001; Hindle et al., 2002]. In such a setting, how has magma risen to the surface? Have the volcanic edifices been subject to shortening? What has happened at depth?

[5] In the Andes, there have been many descriptions of volcanoes in close proximity to faults. Strike-slip faults, and lineaments that are transverse to the orogen, would appear to be the most common [Marrett and Emerman, 1992; Boudesseul, 1997; Matteini et al., 2002a, 2002b; Petrovich and La Torre et al., 2006]. However, some Andean volcanoes are close to thrust faults. A good example is El Reventador volcano in Ecuador, which formed in a compressional or transpressional setting [Tibaldi, 2005]. Other examples along the Andean Cordillera are Guagua Pichincha in Ecuador [Legrand et al., 2002], Taapaca in Chile [Clavero et al., 2004], Socompa in Chile [van Wyk de Vries et al., 2001], and Tromen in Argentina [Kozlowski et al., 1997; Marques and Cobbold, 2002, 2006; Galland et al., 2005]. In the main volcanic arc of Neuquén province, Folguera et al. [2004] have described thrusts and folds in basaltic lava flows of Pleistocene age. However, the relationships of volcanoes to nearby thrust faults remain enigmatic and insufficiently explored.

[6] Elsewhere, there is evidence that some intrusions, which formed at depth in a context of crustal thickening and now at the surface as a result of exhumation (see review by Hutton [1997]), have been associated with major thrust faults. Good examples are the Boulder Batholith of Colo-
rado [Kalakay et al., 2001; Lageson et al., 2001] and the Idaho-Bitterroot Batholith in the United States [Foster et al., 2001]. Could such associations be representative of what happens at the roots of some Andean volcanoes?

[7] In this paper, we describe structural and geochronologic evidence that Tromen volcano formed during active compression in an Andean back-arc context. Tromen volcano rises to 4000 m in a thick-skinned fold-and-thrust belt, on the western margin of the Neuquén basin, Argentina (Figure 1). Its volcanic products are unconformable upon Mesozoic strata. The major eastward verging Tromen thrust dips beneath the volcano and bends around its base (Figure 2) [Kozlowski et al., 1996; Marques and Cobbold, 2002]. We describe the relationships between the volcanic products of Tromen and the tectonic structures beneath or within it. To constrain the timing of volcanic and tectonic events, we have obtained new 39Ar-40Ar ages on the volcanic rocks. On the basis of our results, we discuss possible processes governing the rise of magma and the eruptions of Tromen during horizontal compression.

2. Geological Setting

[8] The Neuquén basin lies in NW Patagonia, to the east of the Andean Cordillera (Figure 1). It contains up to 5000 meters of Mesozoic strata. Synrift sediments of the Pre-Cuyo group and organic-rich mudstone of the Los Molles formation accumulated during a phase of late Triassic to early Jurassic extension that heralded the break-up of Gondwana (Figure 2c) [Vergani et al., 1995; Franzese and Spalletti, 2001; Franzese et al., 2003]. Jurassic and Early Cretaceous strata marked a phase of postrift subsidence on the Pacific margin. The Jurassic Lotena formation is mainly of detrital origin, but includes some carbonates (La Manga member). After the evaporitic Auquilo formation came sandstone of the Tordillo formation. The late Jurassic saw a second accumulation of organic-rich mudstone (Vaca Muerta formation). After another sandstone interval (Mulinchinco formation) came a third thickness of organic-rich mudstone (Agrio formation). The evaporitic Huitrén formation then marked the end of marine conditions. The mudstones of the Los Molles, Vaca Muerta and Agrio formations are the main source rocks for hydrocarbons in the Neuquén Basin. During the Late Cretaceous, continental red beds (Neuquén Group) accumulated regionally in a foreland setting [Cobbold and Rossello, 2003]. During the Cenozoic, continental sediment and volcanic rocks accumulated here and there within the basin.

[9] Several periods of tectonic inversion have been recognized in the Neuquén basin. They correlate with the Peruvian, Incaic, and Quechua phases of Andean orogeny [Groebner, 1929; Cobbold and Rossello, 2002, 2003]. The Peruvian phase (Aptian to Campanian) was responsible for E-W shortening and resulted in folds and thrusts, trending N-S, including the Andacollo-Loncopué Fault System (Figure 1). This first phase also resulted in a regional unconformity between the Agrio and Huitrén formations. The foreland basin then filled with continental sediment up to 2500 m thick. During the Incaic phase (Eocene), oblique convergence at the Pacific margin led to right-lateral transpression and strike-slip faulting [Cobbold and Rossello, 2003]. It also may have resulted in the emplacement of a swarm of NE trending bitumen dikes [Cobbold et al., 1999]. Some authors argue that the Incaic phase was followed by an extensional phase of Oligocene to early Miocene age [Jordan et al., 2001; Folguera et al., 2002], but Cobbold and Rossello (2002, 2003) have discussed evidence for compression during these periods. The Quechua phase (Neogene) resulted from more frontal convergence at the Pacific margin. It reactivated previous structures, and formed new ones (such as the Tromen thrust, Figure 1 [Kozlowski et al., 1996; Cobbold and Rossello, 2003]). The Incaic and Quechua phases were both contemporaneous with periods of rapid convergence at the Pacific margin of South America [Pardo-Casas and Molnar, 1987; Somoza, 1998; Cobbold and Rossello, 2003].

[10] Because Neogene deposits are scarce, it is not clear how long compressional deformation has lasted. According to Kozlowski et al. [1996] and Folguera et al. [2002], it ended before 12 Ma. More recently, Folguera et al. [2005, 2006], Ramos and Folguera [2005], Kay et al. [2006a] have argued that compressional deformation stopped during the early Pliocene, and that active tectonics is extensional. However, GPS measurements [Klotz et al., 2001], in situ stress measurements [Guzmán et al., 2005], and structural and geomorphological data [Cobbold and Rossello, 2002, 2003; Folguera et al., 2004, 2005; Galland, 2005; Galland et al., 2005] provide evidence for current horizontal compression.

[11] In Neuquén province, the cordiller and its foreland have a long history of magmatic activity. During the Eocene, a thick pile of andesite lava flows accumulated to the west of the Andacollo-Loncopué Fault System (Figure 1) [Jordan et al., 2001], and shallow level intrusions were emplaced in the Agrio fold belt [Llambías and Rapela, 1988; Cobbold and Rossello, 2003]. During the Oligocene and Miocene, the magmatic arc migrated eastward [Vergara and Munizaga, 1974; Suarez and Empanaran, 1995; Muñoz et al., 2000] to the north of the Cortaderas fault zone (Figure 1), magmatic activity with an arc geochemical signature reached as far as 500 km east of the trench [Kay et al., 2006b; Kay and Copeland, 2006]. It accounted for 1000 m of basaltic lava flows in the Huantraico syncline, 200 km east of the frontal volcanic arc (Figure 1) [Ramos and Barbieri, 1988; Kay and Copeland, 2006], and early activity of Chacahuanén volcano (Figure 1) [Kay et al., 2006a, 2006b]. Then, during the Pliocene and Holocene, magmatism with an arc geochemical signature migrated westward and concentrated close to the modern volcanic front [Vergara and Munizaga, 1974; Hickey et al., 1986; Folguera et al., 2002]. However, east of the main arc, large shield-like volcanoes, such as Auca Mahuida, Payun Mattr and Tromen (Figure 1), have back-arc (intraplate-like) geochemical signatures [Llambías et al., 1982; Saat et al., 1993; Kay et al., 2006a].

[12] Although Tromen is a back-arc volcano, it is close to the current magmatic arc (Figure 1). At 3969 m, the summit of Tromen dominates the western Neuquén basin. Its main
period of activity spanned the Pleistocene and Holocene [Zollner and Amos, 1973; Holmberg, 1975; Llambias et al., 1982]. However, there are strong indications that Tromen has continued to be active in very recent, and even historical times. Its northern crater is well preserved and contains deposits of sulphur [Zollner and Amos, 1973]. The uppermost basaltic lava flows, which erupted from it, have fresh blocky surfaces that are free of vegetation. A historical eruption in 1822 has been reported [Simkin and Siebert, 1994], although we have no independent evidence for it.
Figure 2. (a) Geological map of Tromen volcano (modified after Zollner and Amos [1973] and Holmberg [1975]). Boxes indicate areas of more detailed study (Figures 3b, 3c, and 5). For E-W cross section of volcano, see Figure 15. (b) Map of volcano-sedimentary formations of Tromen area. Red circles and associated numbers refer to samples for dating (Table 1) and ages obtained. (c) Stratigraphic column and map legend.
Between 1751 and 1752, a Jesuit priest, Bernardo Havestadt, went near the twin peaks of Tromen volcano and wrote, “the smoke that they sometimes throw out is so thick, black, and abundant, that, even at noon, it darkens the place, and transforms the day into night” (our translation from the Spanish [Havestadt, 1752]).

The volcanic products of Tromen are unconformable upon a deformed Mesozoic substrate (Figure 2) [Zollner and Amos, 1973; Llambias et al., 1982]. The main structure close to the volcano is the Tromen thrust (Figure 2a) [Holmberg, 1975; Kozlowski et al., 1996].

3. Products of Tromen

We describe the main products of Tromen, and the petrological features of the volcanic rocks that we sampled for 39Ar-40Ar dating.

At the surface, the most voluminous products of Tromen are basaltic lava flows (Basalto formations, Figure 2 [Zollner and Amos, 1973; Holmberg, 1975; Llambias et al., 1982]). However, andesitic to rhyolitic lavas and domes are dominant in some places (Pleistocene Tilhue formation, Figure 2 [Zollner and Amos, 1973; Holmberg, 1975]). An early geochemical analysis of Tromen lavas yielded a calc-alkaline signature [Llambias et al., 1982]. Later studies indicated that the Tromen lavas differ from those of the main volcanic arc: they are more alkaline and have intraplate-like signatures, like the lavas of other back-arc volcanoes in the area [Stern et al., 1990; Saal et al., 1993; Kay et al., 2006a].

There are also several deposits of conglomerate on Tromen (Pichichacaico, Huecú, and Agua Carmonina formations, Figure 2). Mostly they contain pebbles or blocks of volcanic material.

Although Miocene to Pliocene basalts (Basalto I and Basalto II formations [Zollner and Amos, 1973; Holmberg, 1975]) crop out in the general area (Figure 2), there is no direct evidence that they erupted from Tromen. In fact, the Basalto I formation, which extends as much as 100 km to the east of Tromen, appears to be older [Groeber, 1929]. In the Tromen area, the Basalto I formation appears to derive from a partly eroded volcanic center to the south of the Huantraico syncline (Figure 1) [Holmberg, 1975; Ramos and Barbieri, 1988]. Thus we shall consider neither of these formations any further.

3.1. Domes of Andesite, Dacite, or Rhyolite

Traditionally, the andesite, dacite and rhyolite of Tromen have been assigned to the Tilhue formation [Groeber, 1929; Zollner and Amos, 1973; Holmberg, 1975]. These rocks crop out at (1) Cerro Tilhue in the south, (2) Cerro Bayo in the east, and (3) El Paso, a central
pass between the western and eastern flanks of the volcano (Figures 2 and 3). In the field, these rocks form massive domes, 100 to 300 m thick, which are unconformable upon the Mesozoic substrate. The domes have a marked internal flow banding, which is continuous around them, despite strong dip variations (Figure 3a). We interpret these domes to be extrusive. Although they were originally described as extrusive [Groebel, 1929; Zollner and Amos, 1973; Holmberg, 1975], the map legend from Zollner and Amos [1973] has them as intrusive bodies, presumably by mistake.

At Cerro Tilhue, several andesite domes lie along an E-W lineament (Figure 3b). In thin section (samples TIL02, TIL03, TIL06, TIL07; Table 1), the texture is porphyritic, vesicular, and holohyaline to spherulitic (Figure 4a). Phenocrysts are 3 to 5 mm long, and account for 10 to 30 vol % of the rock. They consist mostly of plagioclase (An_{20-30}, oligoclase), clinopyroxene, oxyhornblende and biotite, the latter two being rimmed by opaque minerals.

At Cerro Bayo, the Tilhue formation is composite. The oldest rocks are two conglomeratic beds, overlain by two flow units of andesite lava, 50 m thick (Arroyo Chacaico valley; Figure 3a; see also cross section B-B in Figure 13 in section 5.2). Unconformable upon them are domes of andesite, dacite or rhyolite, 200–300 m thick. The andesite lava flows have a porphyritic to glomeroporphyritic texture (samples CB03-11 and CB03-12). Phenocrysts (20 to 40 vol %), in a microlitic mesostasis, are mostly of plagioclase (An_{45}, up to several cm in length), rich in inclusions. A few phenocrysts are of clinopyroxene. The two main domes, including Cerro Bayo, lie along an E-W lineament (Figure 3c). The texture is porphyritic, banded and vesicular. Phenocrysts, up to 5 mm long, are in a holohyaline to spherulitic groundmass (Figure 4b). In the andesite domes (samples CB03-01, CB03-04 and CB03-14; Table 1), the phenocryst content is 20 to 40 vol %; the main phase is plagioclase (An_{20-30}), with minor oxyhornblende and biotite, both rimmed by opaque minerals. In the dacites and rhyolites (samples CB03-05, CB03-06 and CB03-13; Table 1), the phenocryst content is 5 to 10 vol %. They are mainly of plagioclase (An_{25}) and alkali feldspar (sanidine or anorthoclase).

At El Paso, andesitic to rhyodacitic domes are unconformable upon Mesozoic strata. Several domes are overlain by younger flows of basaltic lava. The texture of the rocks (samples ABL05, ABL08 and ABL03-03; Table 1) is porphyritic to glomeroporphyritic, mostly vesicular and banded. The groundmass is holohyaline, sometimes fine-grained (Figure 4c). Phenocrysts (15 to 40 vol %, up to 5 mm long) are mostly of plagioclase (An_{25-50}), with minor clinopyroxene, oxyhornblende and biotite.
<table>
<thead>
<tr>
<th>Rock type</th>
<th>Geological Formation</th>
<th>Locality</th>
<th>Map</th>
<th>Sample</th>
<th>Longitude, deg</th>
<th>Latitude, deg</th>
<th>Altitude, m</th>
<th>Type/Mineral</th>
<th>Age, Ma</th>
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*Formation names are after Zollner and Amos [1973] or Holmberg [1975]. Material dated is whole rock (WR), biotite (Bt), or obsidian fragment. $^{39}$Ar-$^{40}$Ar ages (in Ma) include mean, standard deviation (σ) and information on quality. Plateau ages (PA) have been calculated at 1σ or 2σ level, according to quality of analysis. Pseudoplateau ages (Ps-PA) are for less than 70% of $^{39}$Ar degassing. All errors are displayed at the 1σ level.*
3.2. Pyroclastic Deposits

Pyroclastic deposits, including ignimbrites, have been reported on the western flank of Tromen [Zollner and Amos, 1973]. The ignimbrites contain white lapilli of pumice and obsidian fragments. They crop out widely along the Arroyo Chapúa and lower Arroyo Blanco (Figure 2). In the upper Arroyo Blanco, they fill paleovalleys in Mesozoic strata (outcrop I1, Figure 5). Later paleovalleys, cutting into the same ignimbrite, contain thick conglomerate. According to Zollner and Amos [1973], the pyroclastic deposits correlate with the felsic domes of the Tilhué formation. This is also consistent with the geochemical compositions of the rocks [Kay et al., 2006a]. However, we were not able to establish any continuity or correlation between them. The
source of the ignimbrites is not clear. They could have come from Tromen, or from other volcanic centers in the general area. For example, Payun Matru to the north (Figure 1) [Dessanti, 1973, 1975] has emitted large amounts of pyroclastic material over distances of several hundred kilometers. Nevertheless, for dating, an obsidian pyroclast was selected (sample CM05-01, Table 1).

[23] Direct field evidence for volcanic explosions in the Tilhue formation is relatively rare. At Cerro Tilhué, small outcrops of pyroclastic material are close to the andesite domes. The deposits contain both lithic and andesitic fragments, up to 10 cm wide, within an ashy andesitic matrix. Some of the lithic fragments are rich in oxihornblende and plagioclase and have a fine-grained texture (sample TIL04; Table 1). Near the domes of Cerro Bayo, some andesite flows are strongly banded breccias. Layers of andesite, a few centimeters thick and containing plagioclase phenocrysts in a holohyaline groundmass, alternate with discontinuous ashy layers. We take this association as evidence for partly explosive activity.

3.3. Sills and Dikes

[24] Immediately to the east of the Tromen thrust, an E-W dike, about 1 m wide, links four volcanic necks, three small ones and one large one (Chihuido de Tril; Figure 2). Zollner and Amos [1973] and Holmberg [1975] included these rocks in the Basalto III formation. The smaller necks have annular rims, containing flattened and vertically stretched vesicles, as well as radial polygonal jointing. The large neck now forms an eroded pinnacle, about 50 m wide and reaching a height of 1432 m. The core is of dolerite. In thin section (sample CT05-01, Table 1) it contains rare phenocrysts of zoned plagioclase (An<sub>50</sub>–60), olivine, or clinopyroxene, up to 2 mm in size. The groundmass is a network of plagioclase laths (up to 1 mm long and 0.25 mm wide), between microcrysts of euhedral olivine (partly replaced by clays) and clinopyroxene. Opaque minerals and acicularapatite are the main accessory phases. The nonequant minerals form a strong linear fabric, which is vertical, presumably as a result of extrusion. The annular rim of the Chihuido de Tril is finer grained and displays flow banding, a strong vertical lineation, radial polygonal cooling joints, and a few radial dikes. Adhering to the sides of the eroded neck are remnants of metamorphosed Mesozoic strata (Agrio, Huitrín and Rayoso formations).

[25] As mentioned by Groeber [1929], sills and subvertical dikes crop out in the eroded center and western flank of Tromen (upper Arroyo Blanco valley; Figure 5). Andesite sills (sample ABL06; Table 1) are parallel to folded Mesozoic strata (Mulichinco and Vaca Muerta formations). The texture is holohyaline and porphyritic. Phenocrysts (30 vol %) are mostly of plagioclase (An<sub>32</sub>), or of altered hornblende or biotite. Dikes up to 20 m thick strike almost E-W (Figure 5). They crosscut andesite sills, Mesozoic strata, and andesite domes of the Tilhué formation (Figure 5a). Late conglomerate is unconformable upon the eroded dikes and the ignimbrites. The dikes consist of andesite (samples ABL04, ABL07 and ABL03-09; Table 1). The texture is porphyritic, locally banded. The groundmass is microlicitic (Figure 4d), locally holohyaline or spherulitic. Phenocrysts, up to 5 mm long, are of plagioclase (An<sub>40</sub>–50) or altered clinopyroxene.
3.4. Basaltic Lava Flows

[26] Stacked lava flows of basalt or basaltic andesite form a cover to the entire edifice (Figure 2) [Zollner and Amos, 1973; Holmberg, 1975]. The thickness of each flow is no more than a few meters and surprisingly uniform over large distances. Seven basaltic lava flows were selected for dating (see Table 1). The texture is porphyritic, coarsely vesicular, and banded. The groundmass is microlitic to trachytic (Figure 4e). Phenocrysts (less than 20 vol %) are 5 to 8 mm long. They are mostly of plagioclase (An40 – 70) or olivine, with minor clinopyroxene.

[27] Zollner and Amos [1973] and Holmberg [1975] grouped the basalts into various formations, according to field relationships. The oldest (Basalto III, Pleistocene) is mostly on the northern and western flanks of the volcano (Los Barros and Cerro Wayle areas), and locally crops out on the eastern flank (Los Chihuidos de Tril; Figure 2). The Basalto IV formation crops out locally on the northern and eastern flanks. Forming the main cones of Tromen, Cerro Negro del Tromen and Cerro Wayle are flows of the Basalto V formation. They overlie andesite domes of the Tilhue formation (Figure 5) and some eroded E-W dikes (in the upper Arroyo Blanco valley). Near the summit of Tromen are Holocene basalts and basaltic andesites of the Basalto VI and Basalto VII formations [Llambias et al., 1982].

[28] The basalts appear to be very homogeneous, both in the field and under the microscope. From petrological features alone, it is difficult to assign any one sample to a given formation. Also it is difficult to correlate the formations on the basis of field relationships.

3.5. Conglomerate

[29] On the eastern flank of Tromen, conglomerate of the Pichichacaico formation crops out at altitudes of 1500 to 1900 m, over a distance of about 20 km (Figure 2). Holmberg [1975] considered it to be Pleistocene (older than the Basalto III formation, but younger than the Tilhue formation). It contains rounded pebbles, about 10 cm in diameter, and angular fragments, mostly of basaltic andesite [Holmberg, 1975]. The matrix is a sand of the same composition. The pebbles are of two distinct kinds, according to phenocryst length (5 mm for TR01; 2 mm for TR02; Table 1). Both have a porphyritic vesicular texture, a microlitic groundmass (Figure 4f) and phenocrysts of plagioclase (An30 – 40) and clinopyroxene. In the field, conglomerate of the Pichichacaico formation appears to be overlain by lava flows of the Basalto IV formation (Figure 2a).

[30] Also on the eastern flank of Tromen, the Agua Carmonina formation is a conglomerate, containing upper layers of ash and travertine, whereas the Arroyo Huecú formation is a conglomerate only. All contain rounded pebbles, from 10 cm to 1 m in diameter, basaltic sand and lapilli [Holmberg, 1975]. Because of their greater variety of pebbles, the Agua Carmonina and Arroyo Huecú formations are thought to be younger than the Pichichacaico formation (Figure 2) [Holmberg, 1975].

[31] On the eastern flank of the volcano, another conglomerate, a few meters thick, has an upper crust of calcrite ("tosca") and is unconformable upon underlying strata [Holmberg, 1975]. It contains rounded blocks of basalt and andesite, up to 50 cm wide. We were not able to correlate this material with any other conglomerate in the area.

[32] On the western flank of Tromen (Arroyo Blanco), a thick conglomerate fills paleovalleys in underlying ignimbrite. It contains pebbles of basalt and andesite, up to a few tens of cm wide. In the lower part of the valley, the conglomerate rests directly on Mesozoic strata. It also covers andesite domes (El Paso) and E-W dikes. Apparently overlying it are lava flows of the Basalto V formation. We cannot correlate this conglomerate with the one on the eastern flank of the volcano.

[33] Zollner and Amos [1973] suggested that some, if not all, of these conglomeratic deposits formed during periods of Pleistocene glaciation. Indeed some of the main valleys on Tromen appear to be of glacial origin.

4. New $^{39}$Ar-$^{40}$Ar Ages

4.1. Sampling

[34] Thirty-one samples, from the main volcanic formations or from blocks in conglomerate, were selected for dating (Table 1).

[35] The samples were mainly from lava flows, volcanic domes and dikes. Exceptions were (1) a fine-grained lithic block (TIL04) from a pyroclastic deposit near one of the domes of Cerro Tilhue (see "Products of Tromen"), (2) an obsidian fragment from an ignimbrite (CM05-01) and (3) two volcanic blocks from conglomerate of the Pichichacaico formation (TR01 and TR02; Table 1). Weathered and hydrothermally altered rocks were discarded. We did not sample the youngest basaltic lava flows (Basalto VI and Basalto VII formations) or the oldest ones (Basalto I and II, which did not obviously erupt from Tromen).

4.2. Analytical Procedure

[36] Samples were analyzed with an $^{39}$Ar-$^{40}$Ar laser probe (CO$_2$ Synrad$^\text{TM}$). The analyses were mostly on single whole rock fragments of basalt or andesite (Table 1). Others were on biotite from andesites, dacites and rhyolites and an obsidian poylest from the ignimbrite.

[37] Whole rock fragments and minerals were carefully handpicked under a binocular microscope from crushed rocks (0.3–2 mm fraction). The samples were wrapped in Al foil to form packets (11 mm × 11 mm × 0.5 mm). These packets were stacked up to form a pile, within which packets of flux monitors were inserted every 8 to 10 samples. The stack, put in an irradiation can, was irradiated for 5 hours at the McMaster reactor (Hamilton, Canada) with a total flux of 6.4 × 10$^{17}$ radioactive shocks per square centimeter. The irradiation standard was the Alder Creek Rhyolite sanidine, ACs-2 (1.194 Ma according to Renne et al. [1998]). The sample arrangement allowed us to monitor the flux gradient with a precision of ±0.2%.
Figure 6
The 39Ar-40Ar age spectra of 31 samples of volcanic rocks from Tromen. The error bars for each temperature step are at the 1σ level and do not include errors in the J values. The errors in the J values are included in the plateau age calculations.
between 2050 and 2200 m. At Cerro Tilhué, the earliest dome is at the base of the andesite series at 2031 m, whereas later ones are above 2150 m. Thus, in each area, younger domes stand above older ones, as might be expected for sequential extrusion. However, a few samples from young domes were collected at lower altitudes than those from older domes (examples CB03-14 and CB03-04, Table 1).

For the basalt flows, our ages uphold previous deductions from local field relationships. On the eastern flank of Tromen, samples of the Basalto IV formation are effectively older than the one sample of the Basalto V formation (Table 1), upholding the relationships of Holmberg [1975]. On the western flank, except for ABL03-01, samples of the Basalto III formation are indeed older than those of the Basalto V formation (Table 1), confirming the observations of Zollner and Amos [1973]. However, from one flank to another, the formation ages do not match. This discrepancy probably illustrates the difficulty of correlating over wide areas, especially when the rocks are petrologically very similar.

The ages of the conglomeratic deposits are poorly constrained. Those for pebbles in the Pichichacaico conglomerate are late Pliocene, hence somewhat older than inferred by Holmberg [1975] for the conglomerate itself, but older indeed than the Basalto IV formation.

In summary, our ages show that basaltic and more siliceous products erupted almost simultaneously from Tromen, throughout its 2 Ma of activity.

5. Structure of Tromen

Tromen is at the northern end of the Agrio fold-and-thrust belt (Figure 1). The Mesozoic sedimentary substrate of the volcano is intensely deformed. The main tectonic structure in the area is the eastward verging Tromen thrust, at the eastern foot of the volcano (Figures 2 and 8). According to seismic data, this thrust is deep-seated, involving Paleozoic to Triassic basement [Kozlowski et al., 1996; Zapata et al., 1999]. Jurassic evaporites of the Auquilco formation crop out in its hanging wall, in the overturned Las Yeseras anticline (Figures 2 and 8) [Holmberg, 1975; Kozlowski et al., 1996]. In map view, the Tromen thrust is strongly arcuate (Figure 8). It may have formed that way, under the weight of Tromen [Branquet and van Wyk de Vries, 2001; Marques and Cobbold, 2002], or it could be inherited from basement structures, such as

![Figure 7. The 39Ar-40Ar plateau ages (with errors at the 1σ level) of samples from various geological formations on Tromen. Samples of Tilhué formation are in three groups (from El Paso, Cerro Bayo and Cerro Tilhué). The histogram of plateau ages has been calculated using errors at the 1σ level. Ps-Pa denotes the pseudoplateau age (Table 1).](image-url)
Triassic or Jurassic normal faults [Manceda and Figueroa, 1995; Vergani et al., 1995; Zapata et al., 1999]. Below we list new field observations from the western, northwestern, northeastern, and eastern flanks of Tromen, where erosion has provided the best views of structural relationships (Figure 8).

5.1. Western Flank

[53] Arroyo Blanco is a deeply incised valley on the western flank of Tromen. It provides an almost continuous section through deformed Mesozoic strata and their thin volcanic cover (Figure 9, AA’).

[54] In the western part of the profile, a conglomerate is unconformable upon an eroded westward verging anticline in the Agrio formation (anticline A1, Figure 9). This anticline coincides in position and trend with a topographic escarpment (E1, Figure 9), which can be traced to north and south (Figure 8). A second anticline (A2, Figure 9) is partly masked by pyroclastic deposits. Again, it coincides with a topographic escarpment (E2, Figure 9) that continues to north and south (Figure 8). The pyroclastic deposits are unconformable upon the Mesozoic strata, filling paleovalleys. There are two families of subvertical joints in the ignimbrites (stereogram, Figure 9). Such preferred orientations are unlikely to result from cooling. Instead, the acute bisector probably indicates E/W shortening. The ignimbrites are unconformably overlain by thick conglomerate that also fills paleovalleys (Figure 9).

[55] Further east, oil seeps through a more open anticline (Figure 9) [Cruz et al., 1996]. On its eastern flank are three westward verging reverse faults (F1, F2 and F3, Figure 9). These coincide with a third topographic escarpment (E3, Figure 9) that can be traced to the north (Figure 8), this time in lava flows of the Basalto V formation (Figure 9). The flow units become steeper at the escarpment, but do not change in thickness, and this argues for folding after consolidation, as a result of reverse faulting at depth.

[56] Near the eastern end of the section is a large eastward verging anticline (A3, Figure 9). On its western flank, the Vaca Muerta to Agrio formations are monoclinal. A reverse fault in the Vaca Muerta formation carries striations that are compatible with E-W shortening (stereogram, Figure 9). Andesite sills lie within the Mulichinco and Vaca Muerta formations. To the eastern end of the section, andesite domes of the Tilhué formation are unconformable upon eroded strata of the Vaca Muerta formation (Figure 9).

[57] Subvertical E-W dikes crosscut the sills, the Mesozoic strata, and even the domes (Figure 5). Some of the
Dikes are offset up to 20 m across N-S striking faults, which have reverse and strike-slip components of slip, giving E-W shortening (Figure 10). The easternmost outcrops of Mesozoic strata reach an altitude of 3000 m (Figure 9), some 2000 m above the foot of the volcano and 1000 m below the summit. Thus Tromen has a relatively thin volcanic cover, above a much thicker substrate of deformed sedimentary rock.

Some kilometers north of the Arroyo Blanco valley, pyroclastic deposits, including ignimbrite, crop out along much of the incised valley of Arroyo Chapúa (Figures 2 and 5). They are unconformable on steeply dipping beds of the Agrio formation in the core of a fault-related anticline (Figures 2a and 8). The uppermost layers of redeposited ash (white, Figure 11a) appear to onlap the anticline. In the hinge, a striated fault, left-lateral and reverse, is compatible with NW-SE shortening. Outcrop of pyroclastic deposits (I1) is located on Figure 5a.

**Figure 9.** Structures on western flank of Tromen. (a) Topographic cross section along northern edge of Arroyo Blanco. For location, see Figure 8. Notice three topographic escarpments (E1, E2, and E3). Vertical scale is exaggerated. (b) Geological cross section along Arroyo Blanco. For section line, see Figure 8. Vertical scale is 1.2 times horizontal scale. Stereograms (lower hemisphere, from programs of Angelier [1984, 1990]) show joints in ignimbrite and fault slip data in Vaca Muerta formation. Solid lines are for faults and joint planes, dashed lines are for bedding, dots are for striations, and arrows indicate motion of footwall. For ignimbrite, two families of vertical joints, if they are shear joints, are compatible with E-W shortening. For Vaca Muerta formation, stereograms show data in situ (left) and after tilting of bedding to horizontal (right). The latter are compatible with E-W shortening. Notice that anticlines (A1, and A2) and reverse faults (F1, F2, and F3) correlate with topographic escarpments (E1, E2, and E3). Outcrop of pyroclastic deposits (I1) is located on Figure 5a.

5.2. Northwestern Flank

Between Los Barros and Laguna del Tromen, a volcanic plateau is overlain by a series of thin basaltic flows (Basalto III and Basalto V formations; Figures 2 and 12). In this area, three topographic ridges trend NNE-SSW (B1, B2 and B3, Figure 8). These are anticlinal hinges in the lava flows. The flow units maintain nearly constant thicknesses around the hinges and they have dips of 15° or more on the flanks. Flattened vesicles lie parallel to the layering and have been folded as well (Figure 12). Locally, the long axes of the vesicles, marking a flow direction, are oblique to the current topographic slope and also to the current dip directions. We infer that the folds formed after flow and cooling of the lava. Because of their elongate shapes, the anticlines are unlikely to be due to magmatic intrusion. Because they appear to prolong the topographic escarpments (E1, E2 and E3, Figure 8), which themselves overlie...
Figure 10. Dikes in upper reaches of Arroyo Blanco. For location, see Figures 5 and 9. View is to west. Dikes are offset by two major faults. Stereograms (inset, lower hemisphere) show minor faults. Solid lines are for faults and joint planes, dots are for striations, and arrows indicate motion of footwall. In situ data (left) are as measured. Analyzed data (right) yield acceptable solution for stress tensor (from programs of Angelier [1984, 1990]). Calculated axes of principal stress (five-, four- and three-pointed stars) are for $\sigma_1$ (greatest), $\sigma_2$ (intermediate) and $\sigma_3$ (smallest). Data are compatible with E-W shortening (black arrows).

Figure 11. (a) Fault-related anticline in pyroclastic deposits, Arroyo Chapúa. View is to north. For location, see Figure 5. Upper layers of redeposited ash (white) onlap folded ignimbrite (forming cliffs). Fault (dashed trace) offsets base of ignimbrite and underlying strata of Agrio formation (foreground). (b) Close-up view of fault plane in ignimbrite (located in Figure 11a by arrow). Fault surface dips at 87° toward N112°. Striation plunges 66° to south. Sense of slip is left lateral, reverse. This fault is compatible with NW-SE shortening.
reverse faults in the substrate, we infer that the anticlines are
due to ESE-WNW shortening. Because the folds are gentle,
small-scale fracturing should be enough to accommodate
the small amounts of bending in individual consolidated
layers, and flexural slip may be the mechanism of folding at
the scale of the lava pile.

5.3. Northeastern Flank

[61] To the east of Cerro Bayo dome, sandstones of the
Mulichinco formation crop out widely (Figures 2 and 13),
whereas the underlying black shale of the Vaca Muerta
formation is only exposed in valleys. The bedding dips
almost uniformly at 5–10° to the NE, except across open
folds and reverse faults trending NNW-SSE (Figures 8 and 13).
Holmberg [1975] reported outcrops of the Huitrí formation
in the area. Although his map indicates a hiatus between the
Mulichinco and Huitrí formations, we did not observe one.
Instead, we found the Vaca Muerta formation. Upon it and
the Mulichinco formation, Pleistocene sediment of the
Huecú formation forms a thin unconformable layer
(Figure 13), in which two families of joints define a dihedral
angle that is compatible with E-W shortening (stereogram,
Figure 13).

[62] Along the western part of the cross section (BB’,
Figure 13), in the Arroyo Chacaico (located in Figure 2a),
black shale of the Vaca Muerta formation dips gently to the
NE, beneath the volcanic domes. Further west, alternating
conglomeratic units and andesite lavas of the Tilhué forma-
tion dip at about 20°W and are probably unconformable
upon Mesozoic strata. The inward dip is rather unexpected
for the flank of a volcano. The andesitic domes of Cerro
Bayo are unconformable upon this series (Figure 13). We
therefore suspect that major tectonic or erosional events
occurred during accumulation of the Tilhué formation.

5.4. Eastern Flank

[63] At Las Yeseras (gypsum quarries), evaporites of the
Auquilco formation are in the core of an exhumed anticline,
in the hanging wall of the eastward verging Tromen thrust (Figure 2) [Holmberg, 1975; Kozlowski et al., 1996]. This thrust is blind. At the surface is an asymmetric anticline, whereas at depth the thrust involves Paleozoic basement [Kozlowski et al., 1996; Zapata et al., 1999]. Across the plateau of Pampa Tril, Mesozoic strata of the Agrio and Mulichinco formations dip very gently to the west (eastern end of section CC', Figures 8 and 14). In the hanging wall of the Tromen thrust, strata of the Agrio to Tordillo formations dip steeply to the east, and are locally overturned. Fault slip data and joints within the Mulichinco formation indicate E-W shortening (Figure 14).

[64] For more than 10 km across Las Yeseras, the Auquilco formation consists mainly of gypsum, some 400 m thick [Holmberg, 1975]. However, it has one internal layer of limestone, 10 m thick (Calico member, Figures 2 and 14). Between Las Yeseras and La Vega de la Totora, the Calico member forms an asymmetric anticline. In its core is an older limestone (Manga member of the Lotena formation). To the north, the anticline ends against the eastward verging Yesera thrust. Between the Yesera thrust and a local back thrust is a pop-up, which consists mainly of breccia. Angular blocks, tens of meters wide, derive from material of the Auquilco formation and the Lotena group (Figure 14). We interpret this breccia as a tectonic melange. Notice that to the east of the Yesera thrust, sandstone of the Tordillo formation (Figures 2a and 14) and locally black shale of the Vaca Muerta formation crop out. They fill the core of a smooth syncline (Figure 14).

[65] Further west, bedding in the Lotena group forms the asymmetric Totora anticline (Figure 14). Joints and striated faults in sandstones of the Lotena formation indicate E-W shortening. On the western flank of the anticline, the stratigraphic succession is complete, from the Lotena group to the Vaca Muerta formation.

[66] Toward the western end of the cross section, conglomerate of the Pichichacaico formation is unconformable upon the Vaca Muerta, Tordillo and Auquilco formations, filling a paleovalley (Figure 14). The conglomerate hosts two families of vertical fractures, which crosscut the pebbles. The fractures strike at about N80° and N120°, defining a dihedral angle that is compatible with E-W shortening. Both pebbles and matrix show signs of hydrothermal alteration.

[67] To the west of the conglomerate is another breccia, containing blocks, up to 10 m wide, of material from the Auquilco, Tordillo and Vaca Muerta formations (Figure 14). We could not see the nature of the contact between breccia and conglomerate. However, one of the blocks in the breccia derives from a conglomerate, which is very similar to the Pichichacaico formation (photo, Figure 14). We interpret the breccia to be a tectonic melange in a major thrust zone (Totora thrust, Figure 14).

6. Implications for Tromen

[68] According to our 39Ar-40Ar data (Figure 7), the magmatic products and conglomerates of Tromen are of late Pliocene to Holocene age. In addition, some of the geological formations, as defined by Zollner and Amos [1973] and Holmberg [1975], are composite, while others are diachronous. Rather than set up a new classification, we will describe the growth and structure of Tromen, emphasizing the close relationships between magmatism and deformation.

6.1. Growth of Tromen

[69] The 39Ar-40Ar ages for Tromen are consistent with those of other back-arc volcanoes in the Neuquén basin, such as Auca Mahuida [Rossello et al., 2002; Kay et al., 2006a] and Payun Matru [Inbar and Risso, 2001]. Volcanic activity on Tromen started more than two million years ago (Figure 7). The 39Ar-40Ar ages span the range between circa 2.3 Ma and circa 0.5 Ma (Figure 7). There is only one age less than 0.5 Ma, but that is not surprising, because we did not sample the youngest volcanic formations (Basaltos VI and VII). Our youngest age (0.04 ± 0.04 Ma, CB03-10) and reports of historical activity [Havestadt, 1752; Simkin and Siebert, 1994] suggest that Tromen volcano is still active.

[70] From a histogram, there seem to have been two main periods of volcanic activity (Figure 7). The first period covers the emplacement of sills, domes and dikes at El Paso, ignimbrite on the western flank of the volcano, and basaltic lava flows on both western and eastern flanks. Some of the
Dikes may have been feeding lava flows. Clearly, felsic rocks (andesite, dacite & rhyolite) appeared together with basaltic rocks, although the former may have been precursors. The basaltic pebbles within the Pichichacaico formation derive from lava of this period, although the age of the conglomerate itself is unconstrained. However, at Cerro Bayo, andesite flows alternated with conglomerates, indicating that significant erosion occurred between eruptions.

The second period of activity at about 1 Ma includes most of the domes at Cerro Bayo and Cerro Tilhue, and one dome at El Paso. During this period, basaltic lavas flowed at Los Barros. Their ages are similar to that of the Cerro Wayle basalts (1.04 ± 0.06 Ma [Kay et al., 2006a]), suggesting that their source may have been Cerro Wayle, rather than Tromen.

The conglomerates of the Huecu and Agua Carmoñina formations, as well as the calcrete, contain volcanic fragments that are more diverse in origin than those of the Pichichacaico formation. Again, the ages of these formations are unconstrained.

We did not date the Basalto VI and VII formations, but they are likely to be younger than 0.5 Ma. Indeed, Kay et al. [2006a] have obtained an Ar/Ar age of 0.175 ± 0.028 Ma. The flows are relatively thin and less voluminous than the domes and flows of the second period.

6.2. Structure of Tromen

Tromen is in a fold-and-thrust belt (Figure 1) [Kozlowski et al., 1996]. Deformed Mesozoic strata crop out in the eroded center of the volcano, at an altitude of about 3000 m, some 2000 m above the surrounding plain (Figure 15). On the flanks of Tromen, volcanic products are relatively thin. On the western flank, basaltic lava flows are less than 100 m thick in the lower reaches of the Arroyo Blanco, and 200 to 300 m thick in its upper reaches. On the eastern flank, domes are about 200 to 300 m thick at Cerro Bayo (Figure 3). More centrally, volcanic products may be up to 1000 m thick around the main crater of Tromen, as at Cerro Wayle. We conclude that more than 80% of the edifice is Mesozoic sedimentary rock (Figure 15) [see Kozlowski et al., 1996; Zapata et al., 1999]. Thus Tromen, like Auca Mahuida to the east [Rossello et al., 2002], is not like a typical shield volcano, resulting from a thick accumulation of lava. Instead, it consists of a thin volcanic cover, upon an uplifted area of sediment. What structurally distinguishes Tromen from Auca Mahuida is that it has been uplifted as a result of compressional deformation.

Within the substrate of Tromen, the main tectonic structures are northerly trending folds and thrusts, which have resulted from E-W shortening (Figure 8). On both western and eastern flanks, these structures verge outward. To the south, their traces curve around the foot of the volcano [Kozlowski et al., 1996; Marques and Cobbold, 2002].

Thus, on the SE flank, the Tromen thrust strikes NNE-SSW; and on the SW flank, all structures strike NNW-SSE. According to our new observations, analogous trends hold in the north. Thus, on the NE flank, folds and reverse faults within the Mulichinco and Vaca Muerta formations trend NW-SE; whereas on the NW flank, folds within basalt lavas, faults within pyroclastic deposits, and topographic escarpments, all strike NNE.

In the Neuquén basin in general, both thick-skinned and thin-skinned thrusts have been reported, on the basis of surface and subsurface data [Kozlowski et al., 1996; Zapata et al., 1999]. Thick-skinned deformation is manifest in the Cordillera del Viento, to the west of Tromen [Kozlowski et al., 1996], and in the Sierra de Reyes, to the east [Zollner and Amos, 1973; Holmberg, 1975]. Thin-skinned detachments are common in (1) evaporite of the Auquilco formation, (2) black shale of the Vaca Muerta formation [Kozlowski et al., 1996] and (3) evaporite of the Huitríñon formation [Zapata et al., 2001].

On the eastern flank of Tromen, offsets on the Tromen, Yeseras, and Totora thrusts are larger than
1000 m and the sheet dip is about 30°. The oldest rocks to crop out, in the hanging wall of the Tromen thrust, are of the middle Jurassic Lotena group. The Tromen thrust can be seen to continue even deeper on seismic profiles [Kozlowski et al., 1996]. Zapata et al. [1999] have interpreted most of the deformation in terms of basement tectonics, involving reactivation of Jurassic normal faults. In contrast, from outcrop data on the western flank, and from a regional seismic section across the southern flank, Kozlowski et al. [1996, Figure 4] inferred a thin-skinned detachment in late Jurassic shale of the Vaca Muerta formation, reaching under the entire volcano. We cannot rule out this possibility. According to our data for the western flank (Figure 9), the offsets on three reverse faults are smaller than 1000 m, and the sheet dip is about 20° to the west, slightly steeper than the surface slope. The Early Cretaceous Agrio formation crops out on the lower slope, whereas the late Jurassic Vaca Muerta formation crops out nearer the center of the volcano. Because older rocks do not crop out on this western flank, the reverse faults may indeed root into a thin-skinned detachment near the base of the Vaca Muerta formation. Alternatively, the reverse faults may be more thick skinned. In our preferred model, Tromen has formed above a thick-skinned pop-up, between eastward verging fore thrusts and westward verging back thrusts (Figure 15). To test this model would require further seismic data.

At outcrop, fault slip data and joints, whether in the substrate or in overlying volcanic formations, are broadly compatible with E-W shortening, due allowance being made for variations around the volcano [Figure 2a, 9, 10, and 12–14]. The dikes at Chihuido de Tril and El Paso, and the lineaments at Cerro Bayo and Cerro Tilhue, are suggestive of hydraulic fracturing, in a context where the least principal stress trends N-S and the greatest stress trends E-W. The dike orientations and the dome alignments are easy to reconcile with our fault slip data, the current state of stress in the Neuquén basin [Guzmán et al., 2005], and the current direction of convergence at the Pacific margin. They may also be indicative of N-S spreading in the hanging wall of the Tromen thrust, under the weight of the volcano [see Johnson, 1970; Marques and Cobbold, 2002, 2006; Galland et al., 2007].

6.3. Age of Deformation

On regional evidence, the Tromen thrust was active in the Neogene, during the Quechua phase of Andean orogeny [Cobbold and Rossello, 2003]. Kozlowski et al. [1996] suggested that compressional deformation in the area stopped at 12 Ma. More recently, Folguera et al. [2004], Ramos and Folguera [2005], and Kay et al. [2006a] estimated an early Pliocene age for the last phases of compressional deformation. However, recent compression would better account for offset Quaternary glacial deposits [Zollner and Amos, 1973] and for active folds and faults [Cobbold and Rossello, 2003].

The volcanic products and conglomerates of Tromen are unconformable upon a strongly deformed sedimentary substrate. The amount of shortening in the volcanic cover, although difficult to estimate, is probably less than 10%. Thus most of the shortening in the substrate must have accumulated before 2.27 ± 0.10 Ma, when volcanic activity started.

According to our observations, deformation in the volcanic cover accumulated throughout the late Pliocene to Holocene.

1. The domes at Cerro Tilhue and Cerro Bayo lie along lineaments trending E-W. If these were fissures that formed by hydraulic fracturing, E-W compression was active between circa 2 Ma and circa 0.8 Ma (Table 1).

2. The dikes at Chihuido de Tril and in the Arroyo Blanco strike almost E-W. If they formed by hydraulic fracturing under E-W compression, this was active between 2 Ma and 1.6 Ma (Table 1).
results in arcuate thrusts. For testing this hypothesis on 2007, Physical modeling has shown that this is possible, so long for tectonic inversion of earlier arcuate normal faults. Finally, possibility is that the Tromen thrust has resulted from have to be deeper than the Auquilco formation. Another restricted area of evaporite or shale. For Tromen, this would due to the weight of the edifice. Such a weight induces a curvature of folds and thrusts around Tromen may be pressional context. 6.4. Relationships Between Tectonics and Volcanism Lavas of basaltic to more felsic compositions have erupted simultaneously or nearly simultaneously from Tromen in several periods since circa 2.3 Ma (Figure 7). Because Tromen has undergone tectonic shortening, between circa 2 Ma and circa 1 Ma and more recently as well, magmatic intrusions and eruptions have been broadly contemporaneous with deformation. Thus Tromen is one of the first examples where coeval volcanism and thrusting have been documented. El Reventador volcano in Ecuador is another example [Tibaldi, 2005], but from a more transpressional context. Marques and Cobbold [2001, 2002, 2006] and Branquet and van Wyk de Vries [2001] have suggested that the curvature of folds and thrusts around Tromen may be due to the weight of the edifice. Such a weight induces a radial stress field [Johnson, 1970], which modifies the regional stress field of purely tectonic origin. However, arcuate thrust belts may also result from other mechanisms [Lickorish et al., 2002]. One possibility is detachment on a restricted area of evaporite or shale. For Tromen, this would have to be deeper than the Auquilco formation. Another possibility is that the Tromen thrust has resulted from tectonic inversion of earlier arcuate normal faults. Finally, a possibility of interest in the Andean context is that detachment has occurred on a weak molten body of magma. Physical modeling has shown that this is possible, so long as the magma does not solidify [Galland et al., 2003, 2006, 2007; Galland, 2005], in which case the configuration results in arcuate thrusts. For testing this hypothesis on Tromen, deep drilling or geophysical surveys might be appropriate methods.

6.5. Consequences for Magma Transport Beneath Tromen, the magmatic plumbing system is not constrained. However, because Tromen grew in a compressional setting, magma is expected to have formed horizontal or gently dipping intrusive bodies or sills [Hubbert and Willis, 1957; Watanabe et al., 1999; Sibson, 2003]. Not all such sills necessarily result from compressional tectonics, but some may do so, even at crustal scale. For example, a seismic profile across the Hercynian thrust belt of SW Spain has revealed a long, thin, gently dipping and highly reflective body, which is probably a mafic sill, emplaced at midcrustal levels [Simancas et al., 2004; Tornos and Casquet, 2005]. So, under compression what are the mechanisms of magma rise through the Earth’s crust to the surface? Some examples of exhumed intrusions indicate what possibly happens at the roots of Tromen. In the Boulder Batholith of Montana, there is good evidence for a close relationship between thrusting and deep intrusion [Kalakay et al., 2001; Lageson et al., 2001]. The main batholith intruded in the hanging wall of the thrust front of the Sevier orogenic belt. The main frontal thrust has an arcuate trace around the batholith, resulting from the emplacement of magma during the formation of the orogenic belt [Lageson et al., 2001]. The batholith is a tabular intrusive body, a few kilometers thick, gently dipping to the west, and subparallel to the thrust [Kalakay et al., 2001]. In the Idaho-Bitterroot Batholith, many intrusions also follow a major thrust [Foster et al., 2001]. Such associations between intrusions and thrusts suggest that during compression, thrust faults could provide a path for magma toward the surface. Physical modeling has also shown that thrusts can control magma transport in a shortening crust [Galland et al., 2003, 2006, 2007; Galland, 2005; Musumeci et al., 2005]. If magma rises along a thrust, it must move horizontally over a large distance, perhaps in the order of kilometers. Under these conditions, Tromen volcano may not lie vertically above its deep source of magma, and we expect significant horizontal transport of magma within the crust.

In the Andes, other volcanoes lying alongside major thrust faults are also associated with arcuate traces of the thrusts, either around the foot of the volcano, or at a slightly bigger scale. Examples are Guagua Pichincha [Legrand et al., 2002] and El Reventador [Tibaldi, 2005] in Ecuador; Taapaca [Clavero et al., 2004] and Socompa [Branquet and van Wyk de Vries, 2001; van Wyk de Vries et al., 2001] in Chile; and Maipo in Argentina [Nullo et al., 1993]. We interpret such associations as possible indicators that these volcanoes have grown during thrusting and we suspect that lateral magma transport along active thrust faults is a common feature within the crust of the Andean volcanic arc.

7. Conclusions In this paper, we have provided structural and geochronological evidence for the growth of a volcano in an active compressional setting. Tromen is a large back-arc
volcano in the foothills of the Neuquén basin of Argentina. Its substrate consists of deformed Mesozoic sedimentary rocks. Volcanic activity on Tromen started at circa 2.3 Ma or earlier and has continued until today. According to our $^{39}$Ar-$^{40}$Ar ages, basaltic lavas erupted almost continuously during this period. The latest of these eruptions appear to have been historical, so that Tromen is technically an active volcano.

The more felsic products of Tromen appear to have formed during two main periods. Most of the domes near the center of the volcano erupted during the first period (circa 2 Ma), whereas most of those on its eastern and southern flanks formed during the second period (circa 1 Ma). These products erupted through fissures trending E-W, which presumably formed by hydraulic fracturing, in response to magmatic pressure and E-W compression.

On Tromen, basaltic rocks form a thin cover, typically a few hundred meters thick, whereas Mesozoic sedimentary rocks account for about 80% of the edifice. The sedimentary strata have been tectonically uplifted in a fold-and-thrust belt. The main Tromen thrust trends broadly N-S, but bends around the base of the volcano. Tromen appears to be at the center of a wide pop-up, between eastward verging fore thrusts and westward verging back thrusts. The former at least are deep seated. All these tectonic structures appear to have resulted from crustal thickening and dominantly E-W shortening.

The volcanic cover of Tromen is unconf ormable upon its Mesozoic substrate. Most of the shortening in the substrate occurred before the beginning of volcanic activity. However, the volcanic rocks of Tromen have also recorded a significant amount of E-W shortening (perhaps as much as 10%). According to our $^{39}$Ar-$^{40}$Ar ages, tectonic deformation was active at circa 2 Ma, and continued after 1 Ma. Hence Tromen has been volcanically active while in a compressional tectonic setting.

Although constraints at depth are cruelly missing, we suggest that the Tromen thrust may have provided a lateral path for magma on its way to the surface. Such lateral transport of magma may be common beneath arc volcanoes growing in close association with active thrusts.

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